



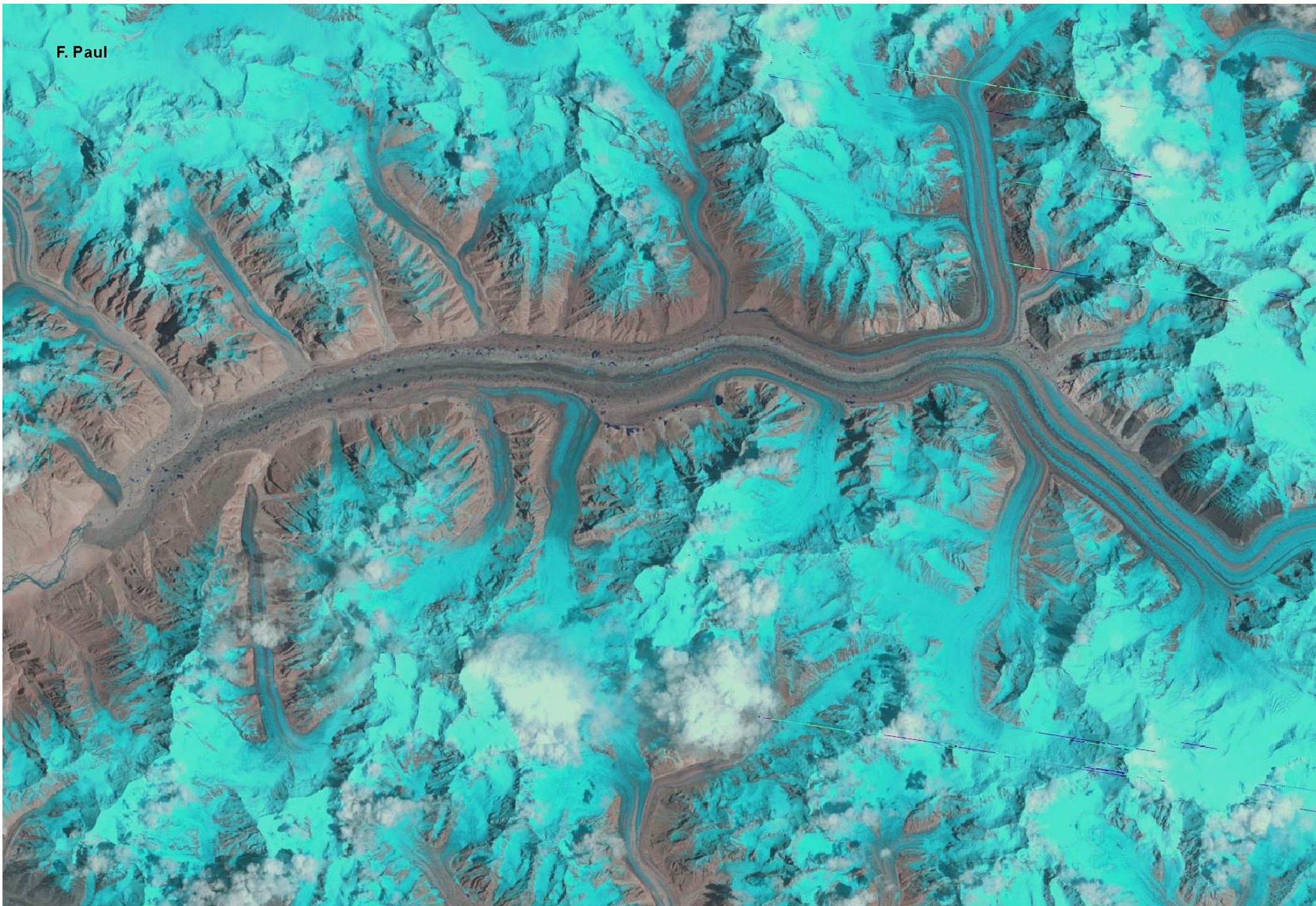
Understanding glacier fluctuations in the Himalaya through simple models

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IISER Pune

Outline:

- * Glaciers (basic processes; importance)
- * Himalayan glaciers (characteristics; data and some puzzles)
- * Results from a simple model glacier
- * ...

F. Paul



Paul, F. (2015)

* Net **accumulation** above end-of-summer **snowline (ELA)**
(snowfall, avalanches, wind-blown snow, ...)

* Conversion of snow into ice
(compaction; reduction in albedo)

* 'slow' **Flow** of ice down-hill

* Ice is a ***hot*** solid => flows as a viscous liquid due to **creep**

$$v \pi R^2 = \frac{\pi R^4}{8\eta} \left(\frac{-\Delta P}{\Delta x} \right)$$

$$R \sim 10^3 \text{ m}$$

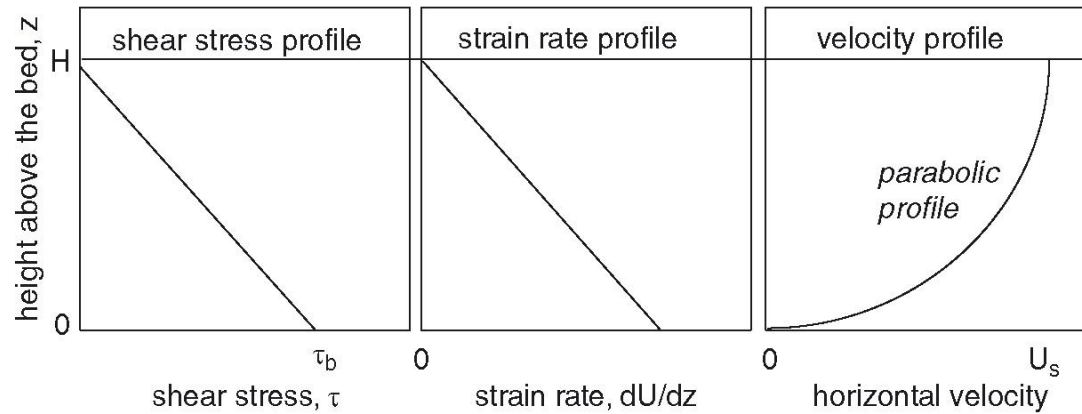
$$v \sim 100 \text{ m} / (10^7 \text{ sec}),$$

$$dP/dx \sim \rho g \sin \alpha = 1000 \times 10 \times 0.1 \text{ Pa/m}$$

$$\Rightarrow \eta \sim \mathbf{10^{13} \text{ Pa-s} \quad !!!}$$

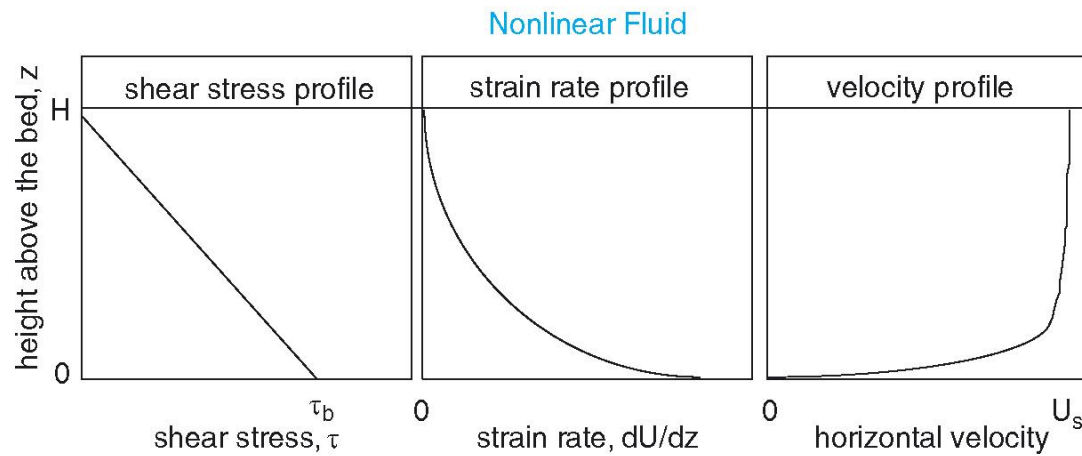
$$(\text{Honey} \sim 10 \text{ Pa-s}, \text{Water} \sim 10^{-3} \text{ Pa-s}, \text{Air} \sim 10^{-5} \text{ Pa-s})$$

Linear or Newtonian Fluid

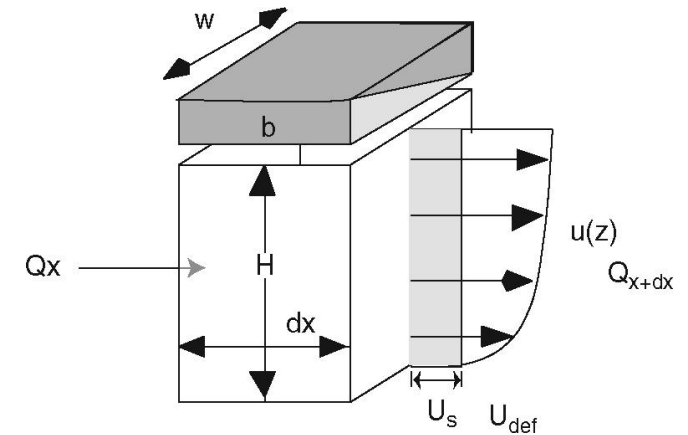


* Viscosity of ice depends on shear stress;
larger shear \Rightarrow lower viscosity
(a **Shear-thinning** liquid)

* there is some **sliding** in addition to the **deformation**



[Anderson & Anderson, 2010]



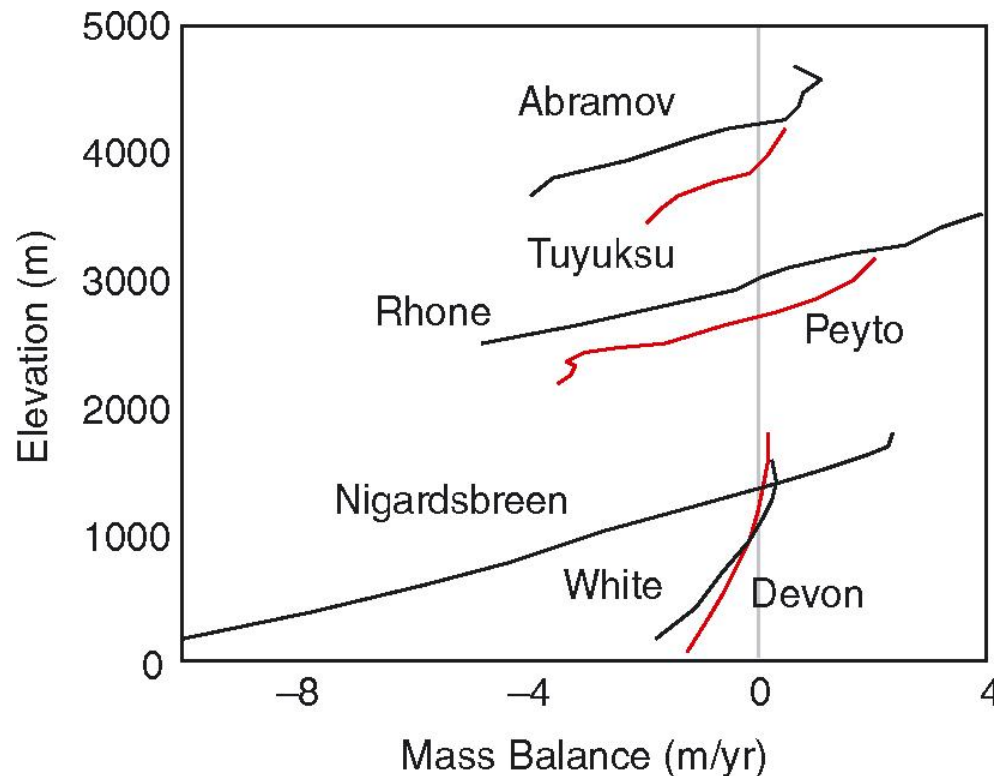
* Once below ELA, glacial ice **melts**.

Local surface **energy balance** determines melt rate

Net radiation budget dominates ($\sim 100\text{W/m}^2$); albedo is important

Latent heat flux, sensible heat flux, ... ($\sim 10\text{W/m}^2$)

* Specific **mass balance** is **linear** in elevation



* ELA is controlled by climate

* Given a steady climate,
glaciers reach a **steady-state**,
net accumulation = net ablation,

$$\int b(z; E) a(z) dz = 0 .$$

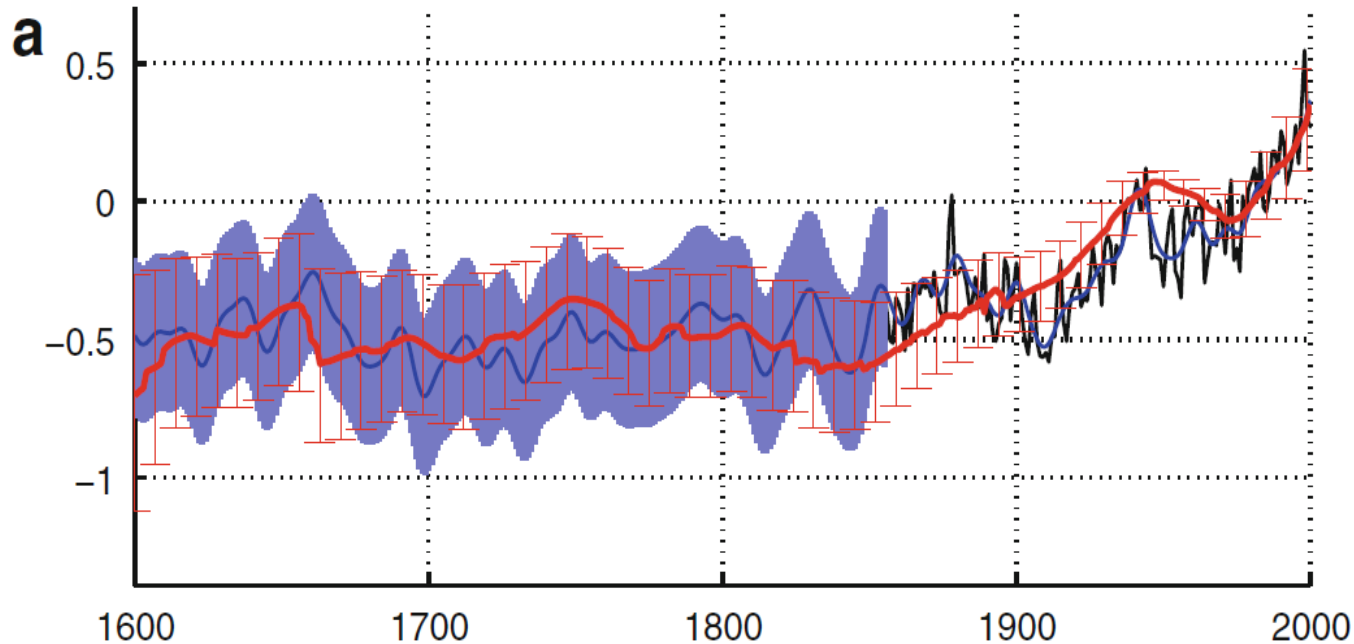
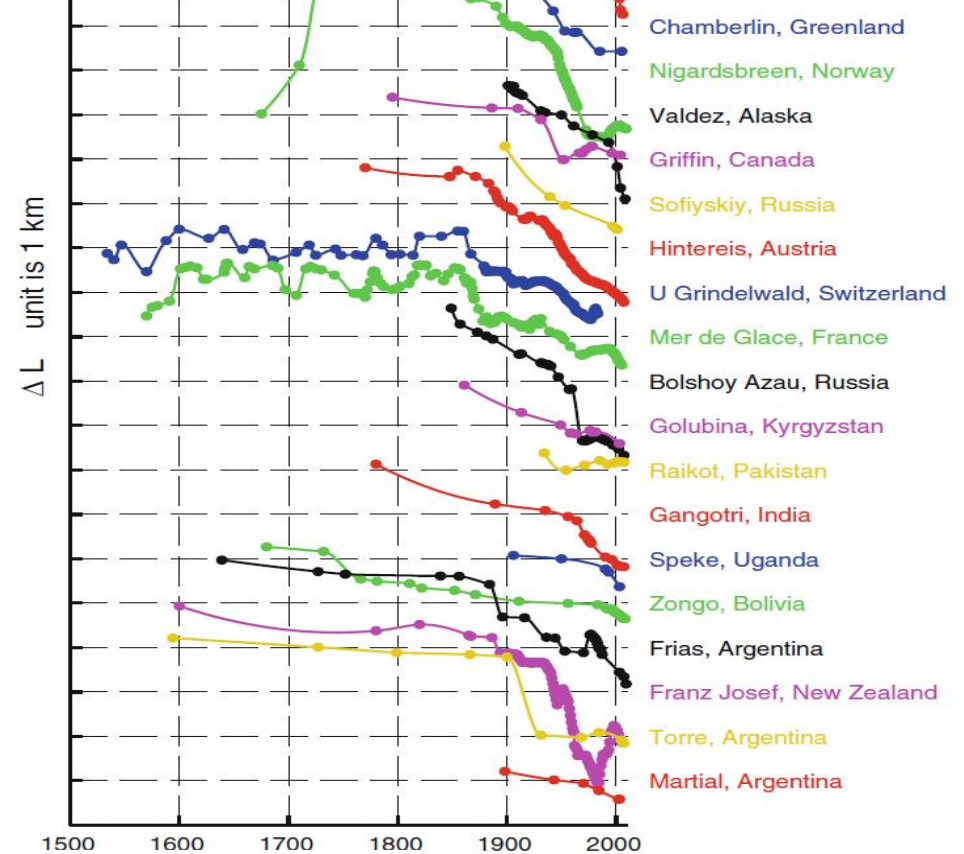
* Response time is large (\sim decades)

Why study glaciers?

□ ** Temperature **reconstruction**

$$\frac{dL'(t)}{dt} = -\frac{1}{\tau}(cT'(t) + L'(t))$$

[Oerlemans, 05]



** In a warming climate glaciers shrink => mean **sea level rises**

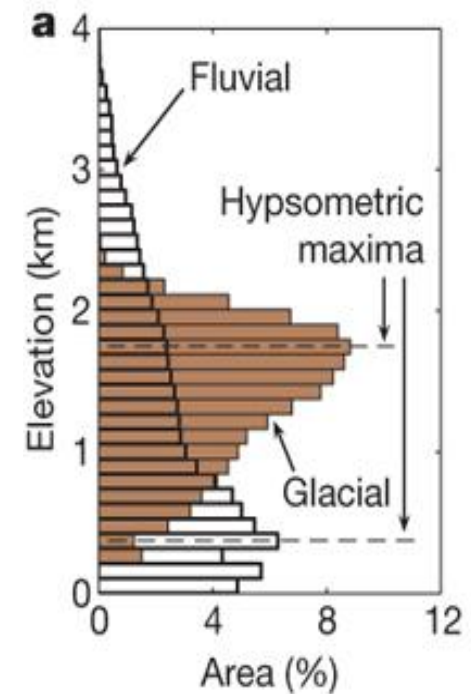
Antarctic Ice Sheet ~ 70m

□ Greenland ice sheet ~ 10m

□ mountain glaciers and ice caps ~ 1m

The present rate of sea level rise ~ 3mm/yr

** Glaciers shape mountain **landscapes** (erosion/transport/deposition)



* Glacial buzzsaw (eg Egholm et al, 2009)

** Glacier **melt-water** feeds rivers:

Glacier/snow melt contribution to present runoff in the Ganga:

- * down-stream regions get a **relatively small** contribution from glaciers
- * more important for the **upstream areas**
- * snow-melt contribution becomes significant during MAM

- * Maurya et al, 2010 ~**10%** in **winter**, ~**30%** in **summer** (at Rishikesh)
- * Bookhagen et al, 2010 ~**18%** (at Rishikesh/upstream areas)
- * Immerzeel et al, 2010 ~**10%** for the whole basin
- * Siderius et al, 2013 ~**1-4%** of annual runoff at Farakka (**12-38%** during MAM)

A **large spread** in these different estimates.

(Better hydro-meteorological data-set; a detailed model with (dynamic) glaciers/snow; better descriptions of **debris covered area** and **avalanches** are needed.)



Himalayan glaciers

- * relief/topography
avalanches
debris cover

ISM, WD, ...

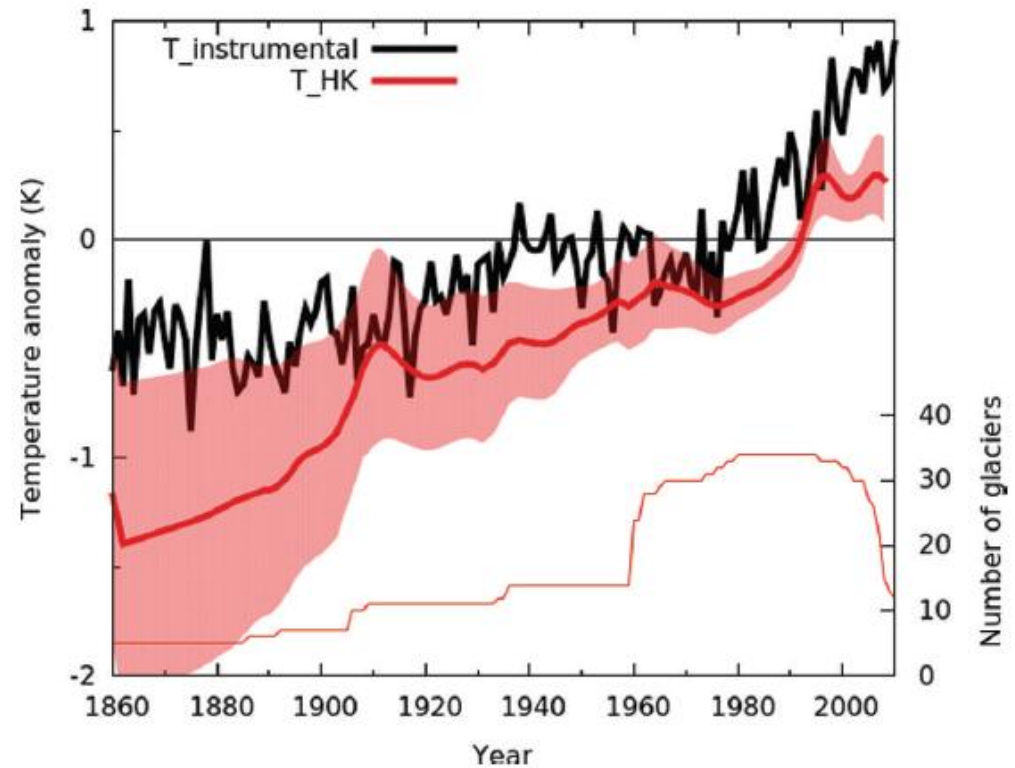
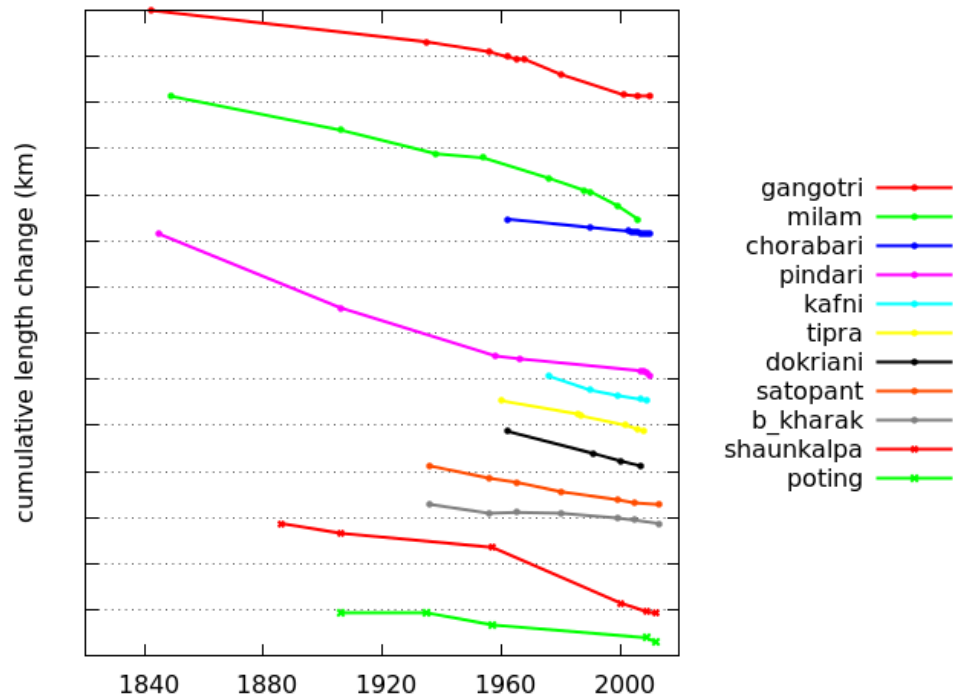
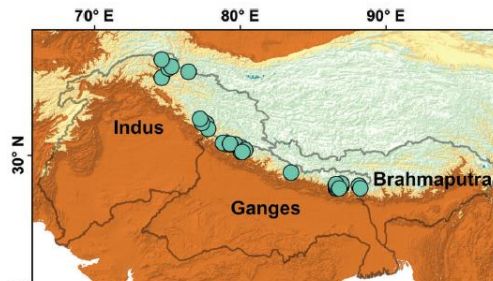
- * data gaps
(Long term data on Himalayan glaciers, Rivers, and climate is missing)

Questions:

- * Recent changes (over decadal/century scale)
- * Future changes under given climate scenarios
- * Contribution of glacier melt-water in river run-off (and its future)
- * Debris effects (on energy balance, flow, hydrology)
- * Bedrock geometry and ice volume
- * Role in landscape evolution, ...

* Present status:

Himalayan glaciers are **retreating/shrinking** more or less steadily over last ~150 yrs



[Banerjee and Azam, 2015]

What is their future?

- * Over the past five decades,
Net sp. Balance ~ **-0.5 m** water equivalent (similar to global trends).

A crude estimate of their lifetime,
Typical mean thickness/ thinning rate ~ 100/0.5 yrs
~ **200 yrs**

(large glaciers are thicker)

- * Another estimate:

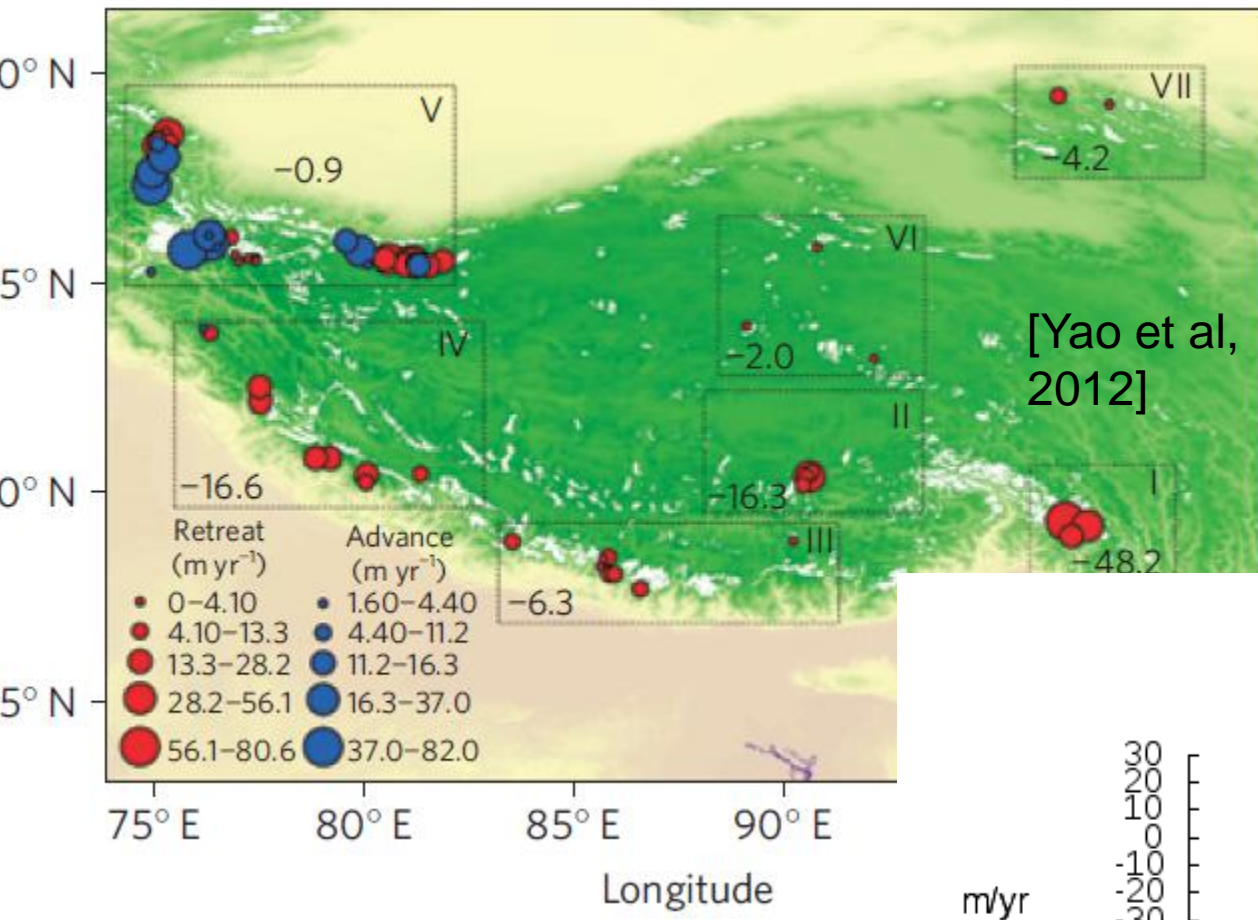
Present ELA in the Himalaya ~ **5000-5500m**

ELA is moving up (due to climatic change) with typical rates of
~10m/yr

*Some accumulation area would be left for the next **few hundreds years** and the higher reaches would remain glacierised.
(available area at higher elevations decreases sharply)

(ice-thickness feedback, ice-albedo feedback, changing glacier area ...)

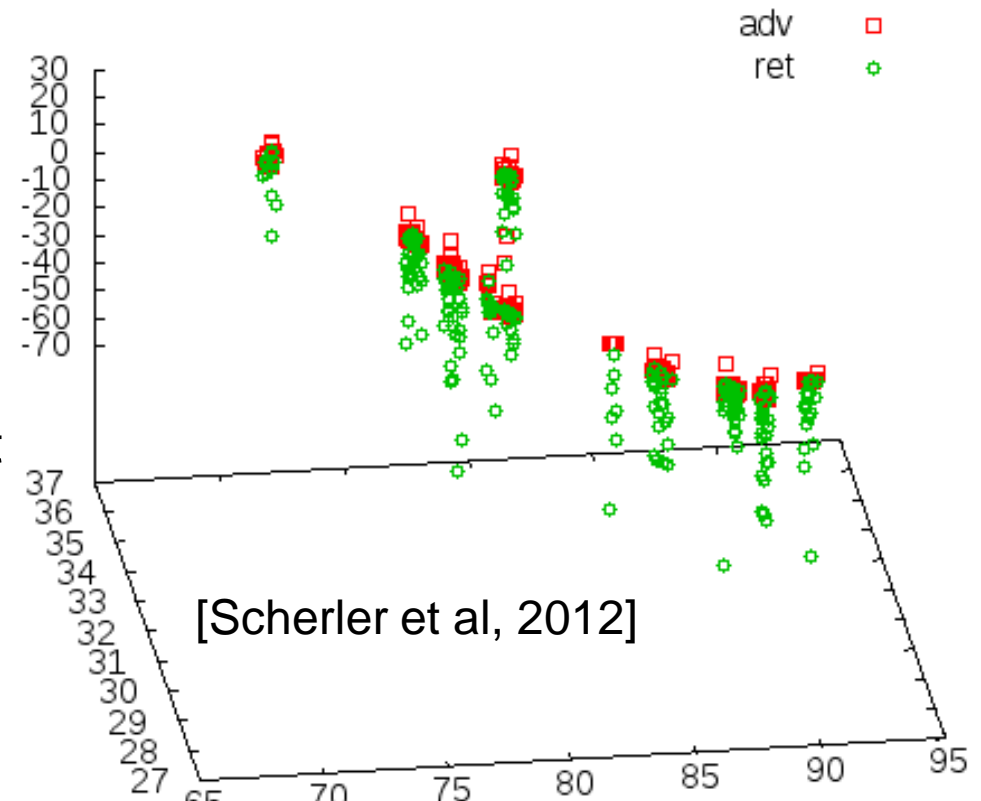
* Data on recent glacier fluctuations in the Himalaya



[Yao et al, 2012]

Puzzle 1:

Wide **variations** in the retreat rates
(many advancing/stationary glaciers!)



Puzzle 2:

In the Himalayan data set,

- 128 debris covered (>10%) glaciers:
 - 52% retreating; 48% advancing/stationary
 - 61 relatively debris free (<10%) glaciers:
 - 82% retreating; 18% stationary/advancing [Schweber et al 2012]
- [Banerjee and Shankar, 2013]

But,

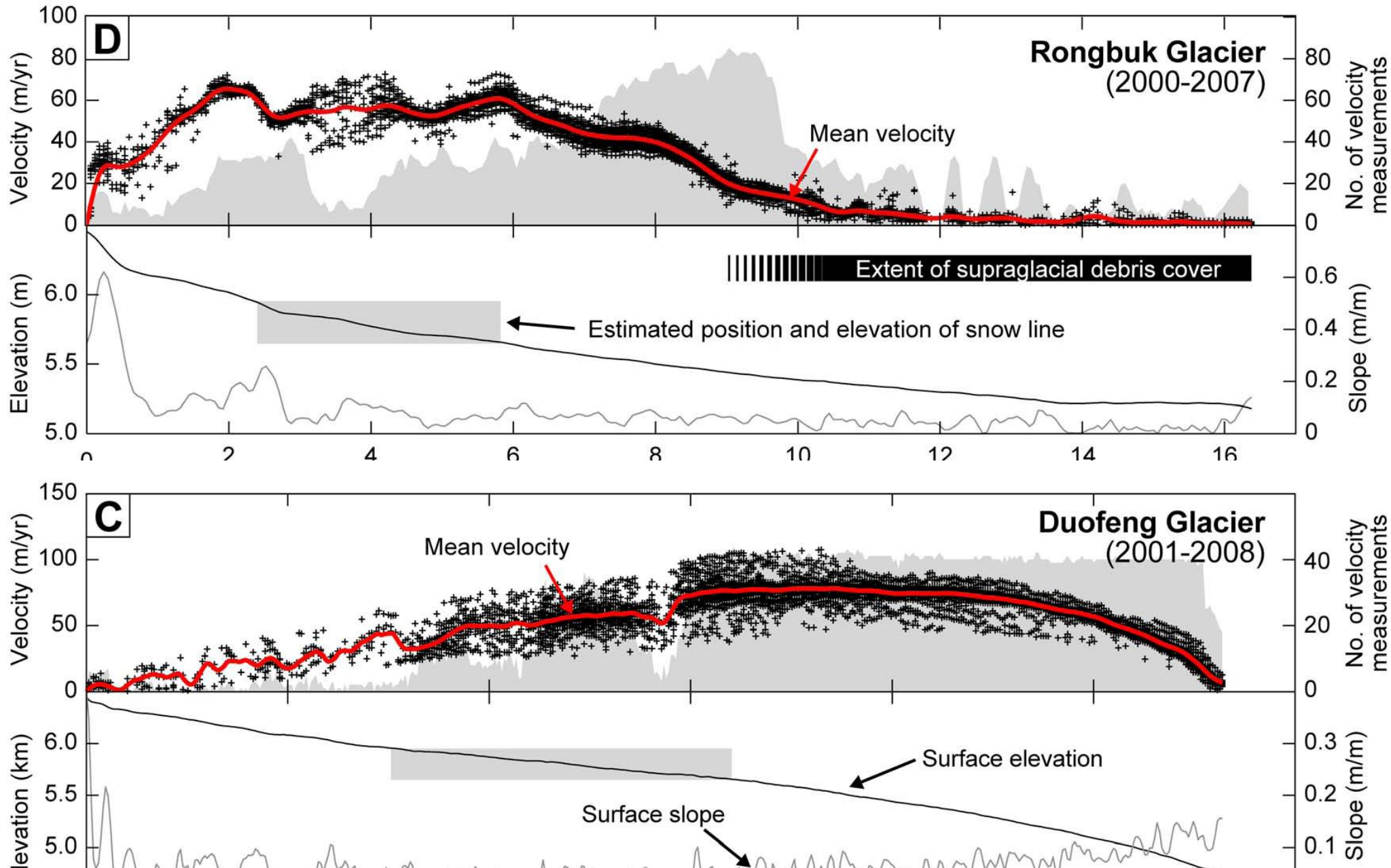
- Both must be seeing similar climate
- **Similar thinning** in clean and debris covered glaciers

[Gardelle et al 2012; Kaab et al 2012; Nuimura, 2012]

(*** Gardelle et al, 2013:: larger/similar/smaller thinning rates in debris covered glacier depending on region, in Pamir-Karamoram-Himalaya)

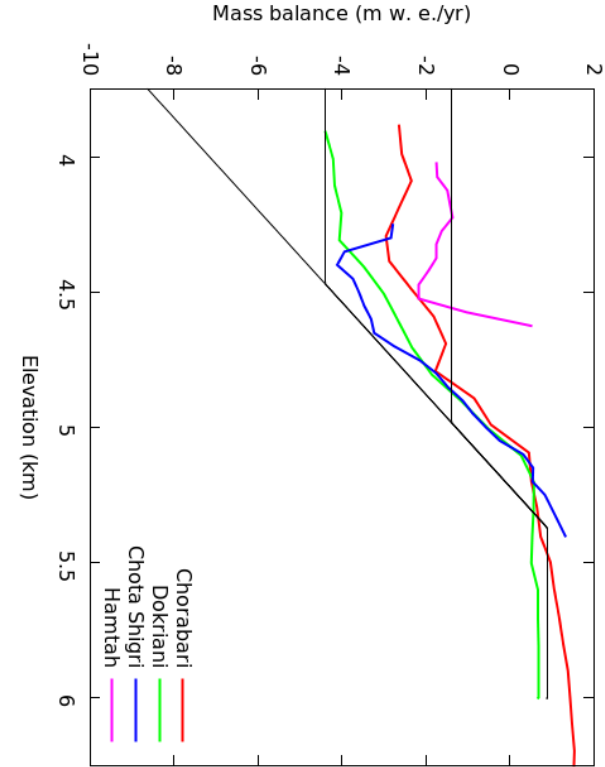
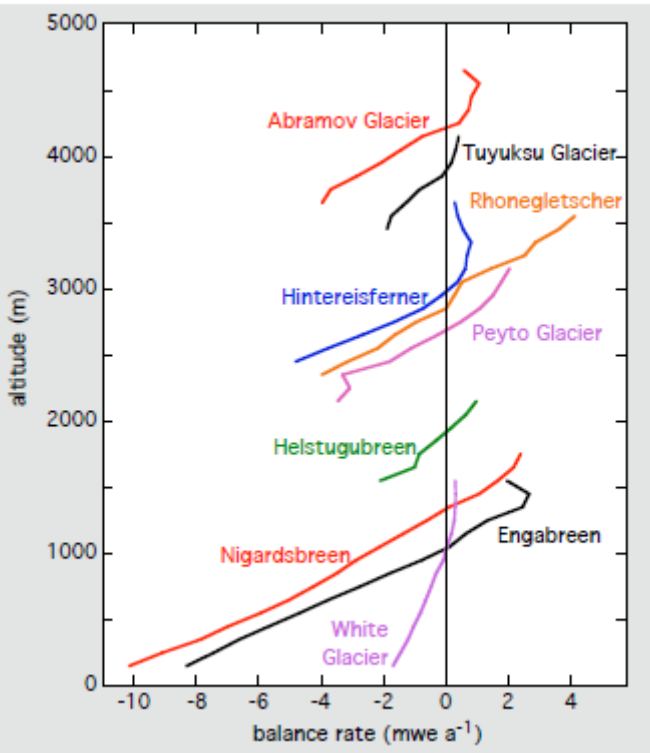
●Puzzle 3:

● Many extensively debris covered glaciers with a low velocity ($v < 2.5 \text{ m/yr}$)
“**stagnant**” tongue and a **steady** front





* A (simple) description of debris-covered glaciers is required



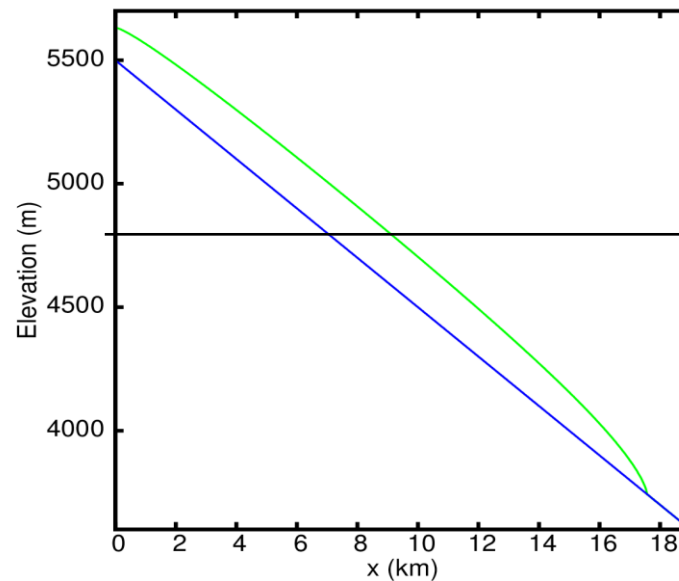
* 1d Flowline model of glaciers:

$$\partial_t H = -\partial_x(HU) + B$$
$$U = f_d(\rho g H \partial_x(H + z))^3 H + f_s(\rho g H \partial_x(H + z))^3 / H$$

H: ice thickness

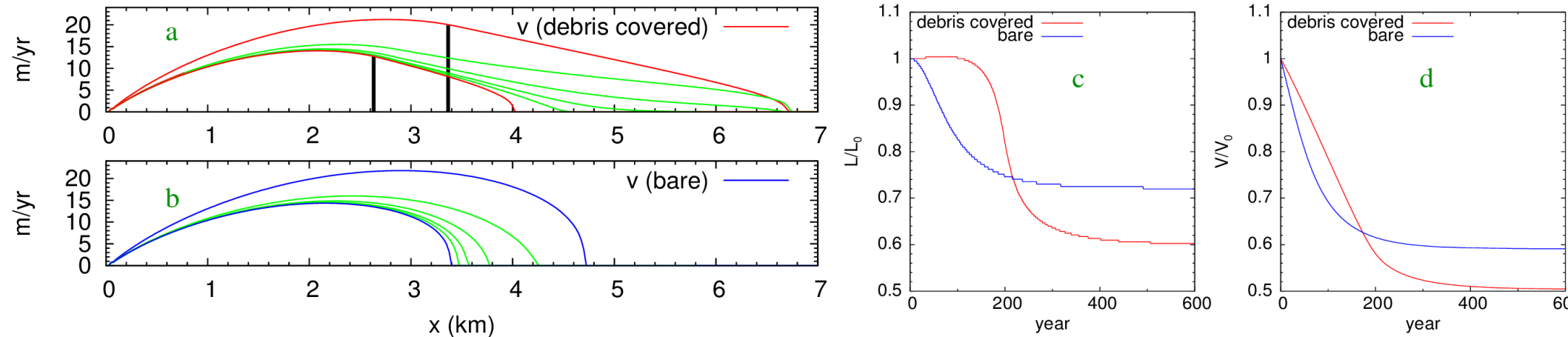
U: flow velocity (function of H, slope)

B: mass balance function (two parameters: balance gradient, ELA)



•Results:

Banerjee and Shankar, 2013

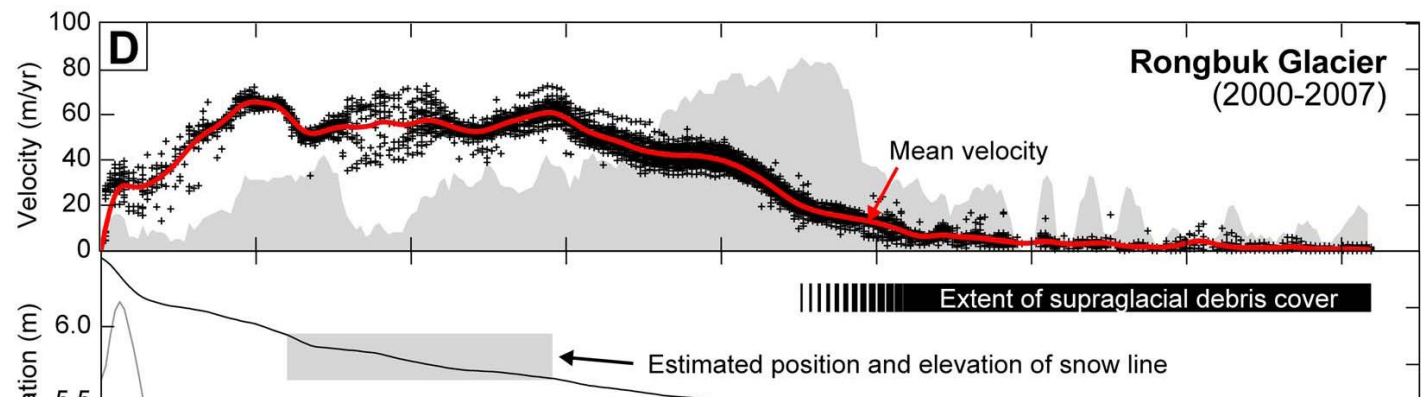


- For debris covered glaciers, initial period of **stagnation** and **thinning** (*downwasting*), without any frontal retreat

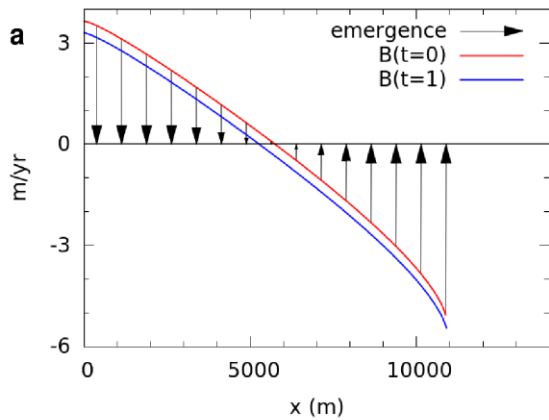
- For clean glaciers, thinning and retreat are simultaneous

- Volume loss curves are similar

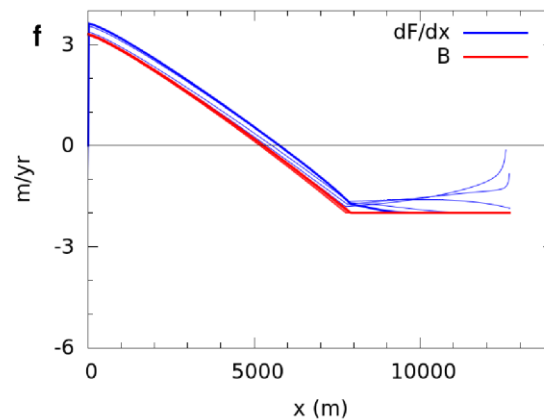
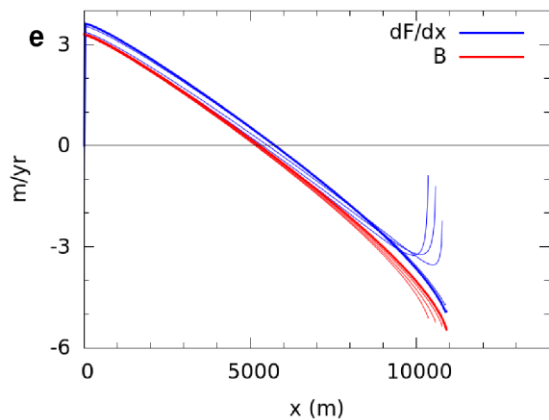
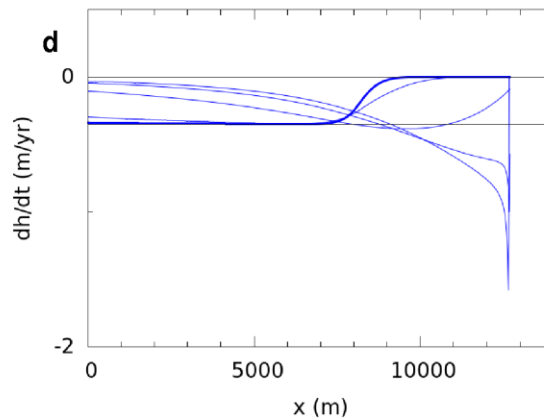
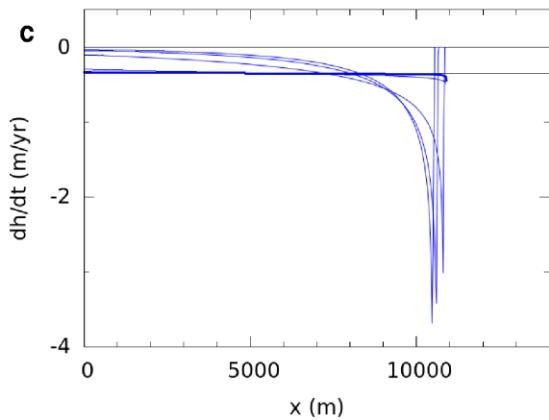
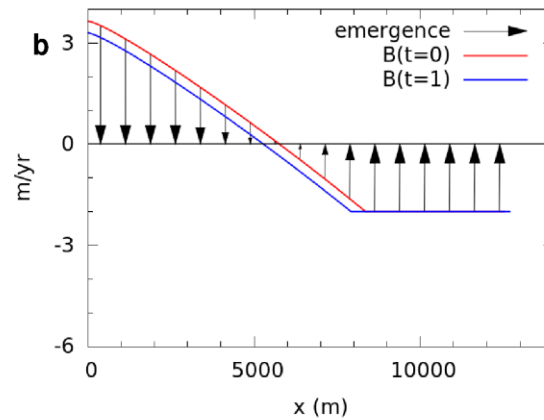
● **Puzzle 3** resolved!



Glacier A

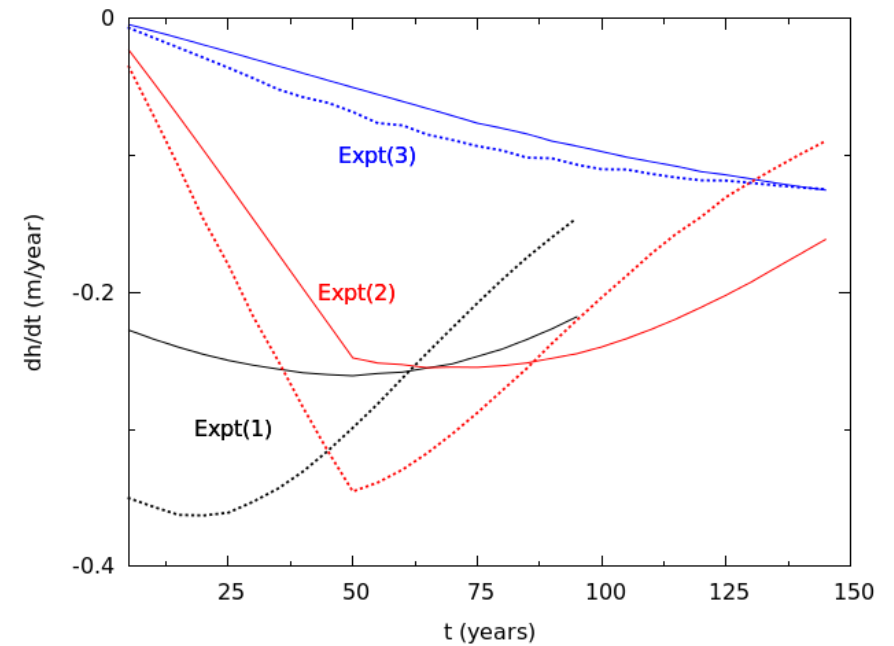


Glacier B



$$\partial_t H = -\partial_x (HU) + B$$

- * the steady state is balance btn emergence (submergence) and melting (accumulation)
- After a warming event,
- * initial volume loss controlled by mass-balance change
- * velocities relaxes slowly
- * subsequently, the evolution of flux changes controls thinning



[Banerjee, 2016*]

● Resolution of **Puzzle 2**:

● Some of the debris covered glaciers showing small or no retreat, are losing ice by thinning (downwasting).

● Assume all debris covered glaciers with $> 5\%$ stagnant area ($v < 2.5\text{m/yr}$) are shrinking,

● The percentage of shrinking debris covered glaciers is up from **52%** to **73%**.

● **$(73 + 27 = 100; 82 + 18 = 100)$**

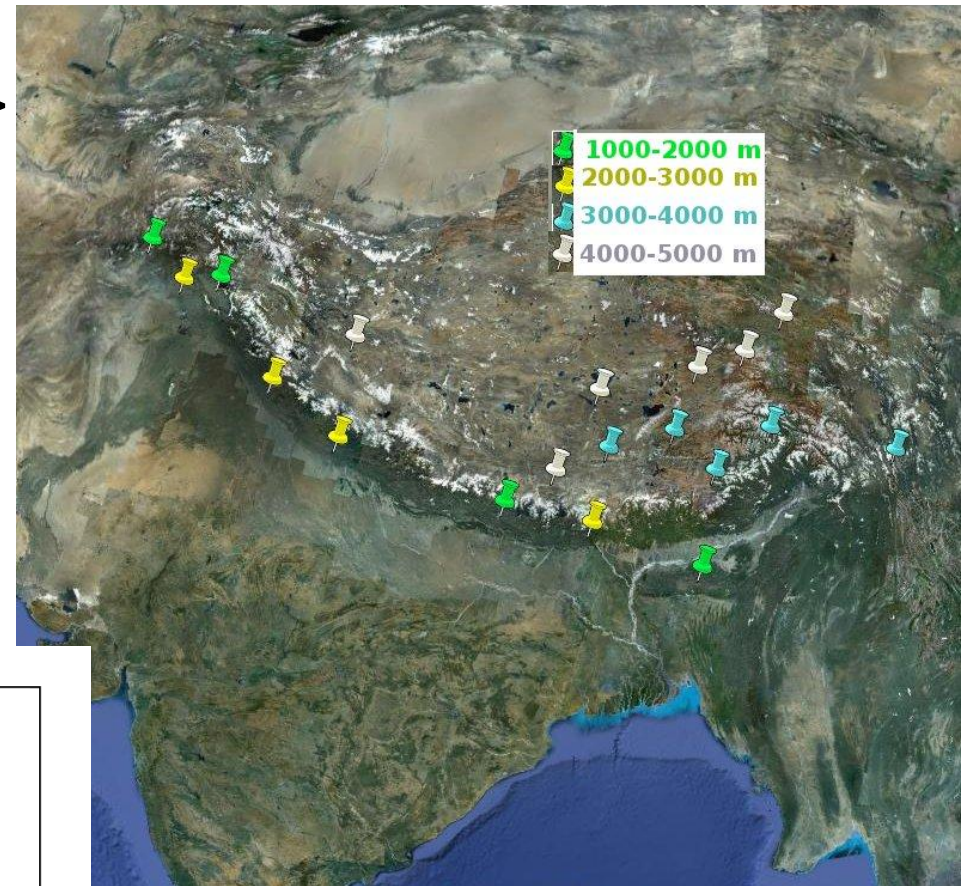
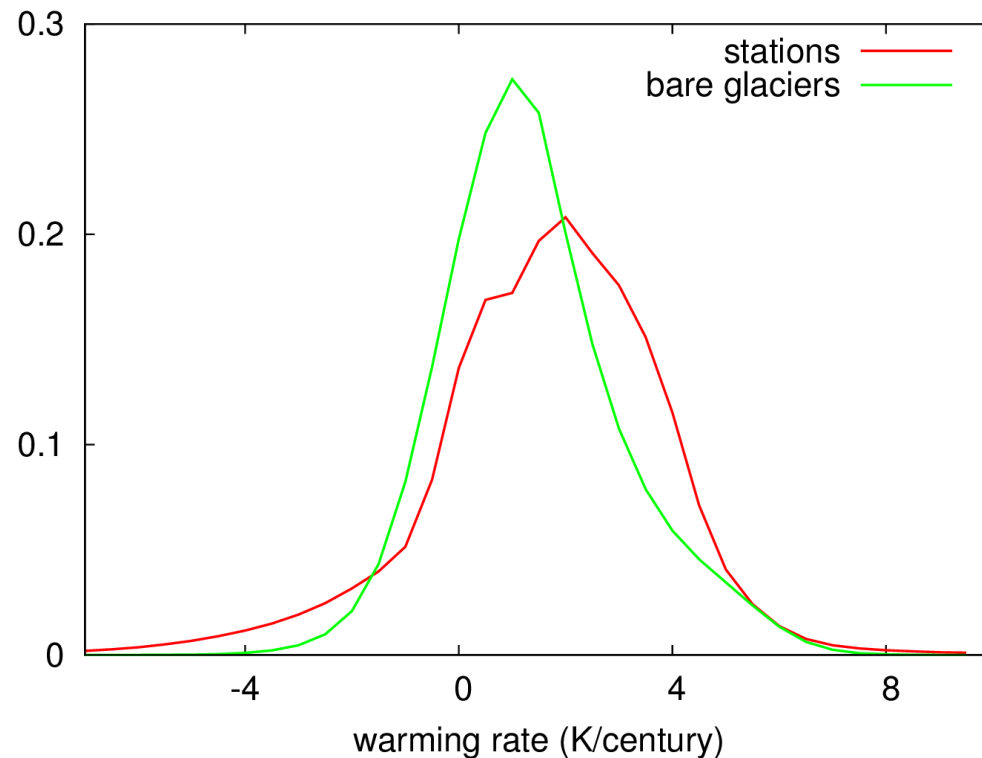
Retreat rate distribution for clean glaciers =>
warming rate distribution

1.4 (1.7) K/century

consistent with available weather station
data

(GHCNM, 19 stations)

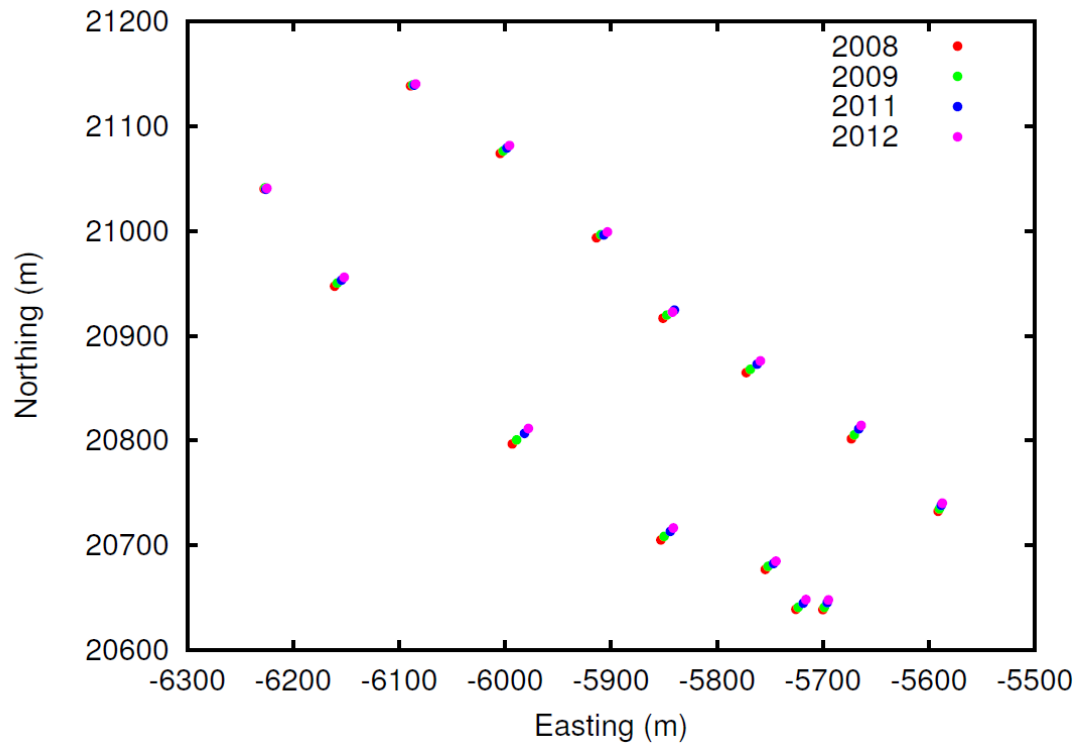
1.6 (2.2) K/century



□ (**Puzzle 1** solved)



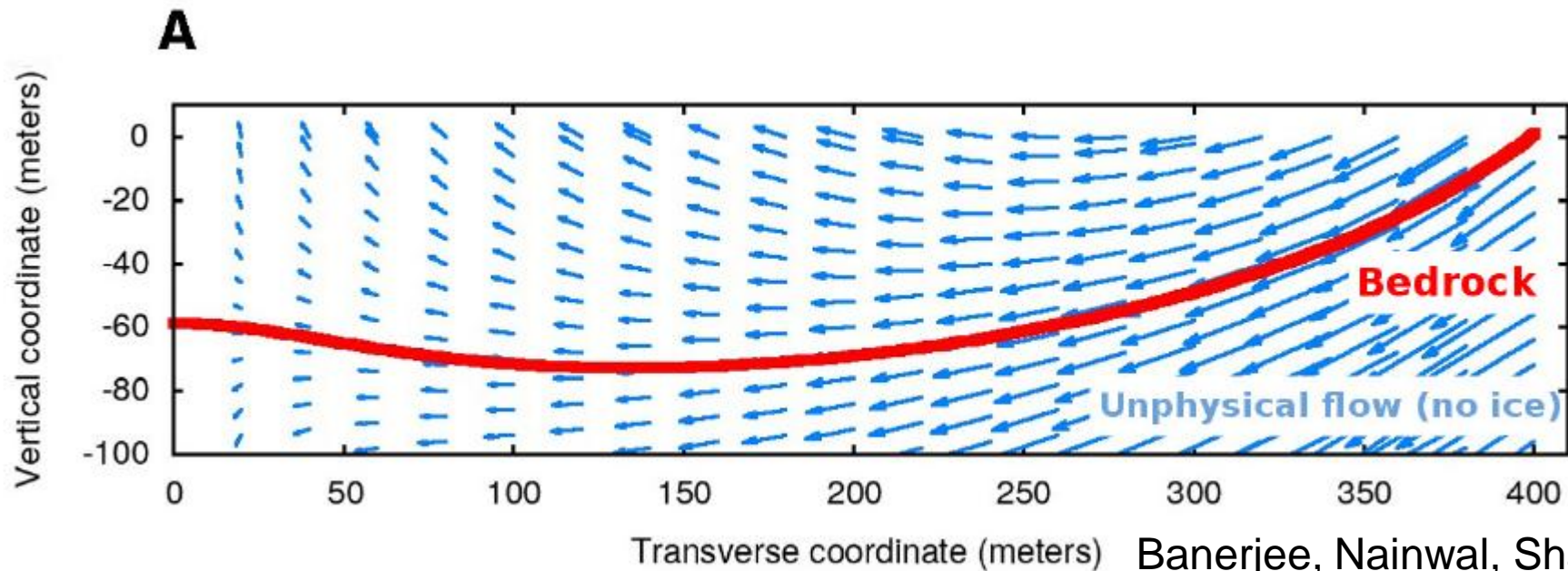
Thank You.



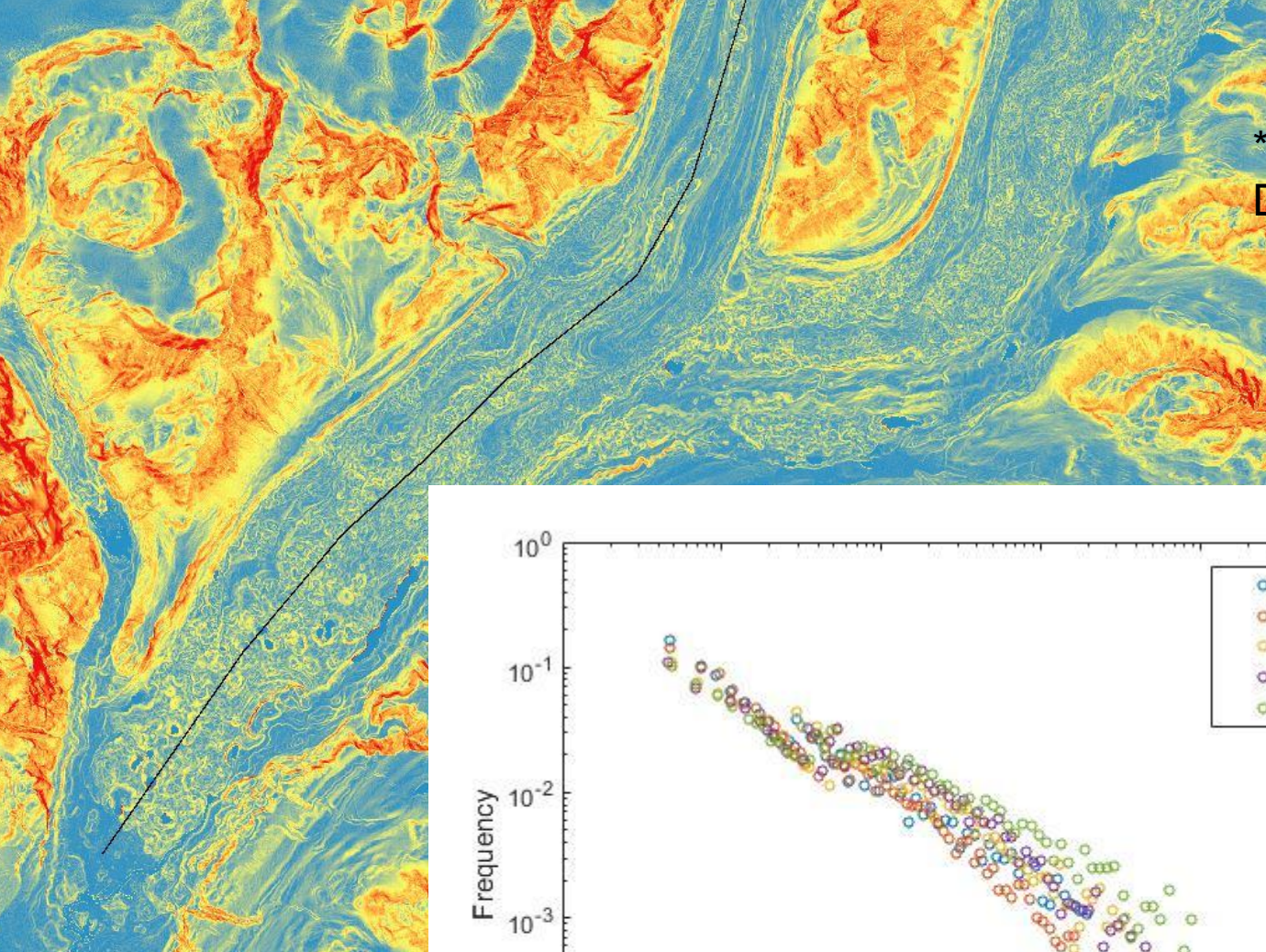
Inversion of surface velocities

$$\partial_j \sigma_{ij} + \rho g_i = 0$$

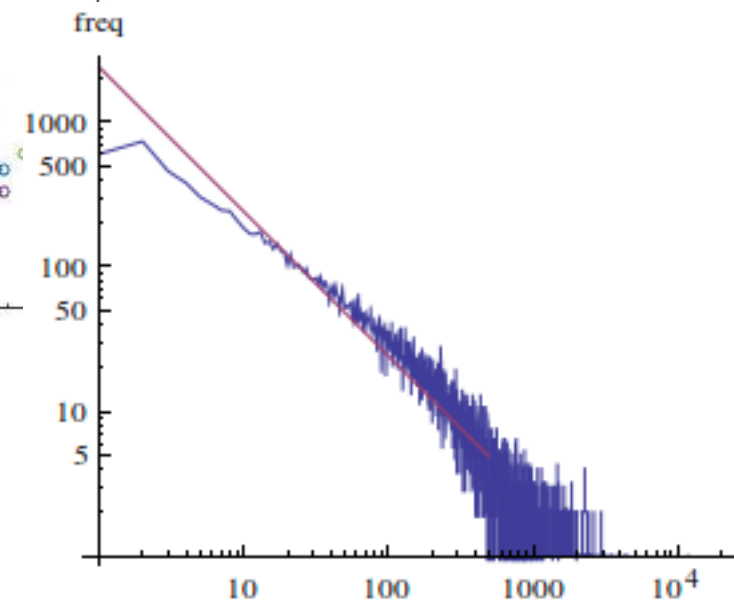
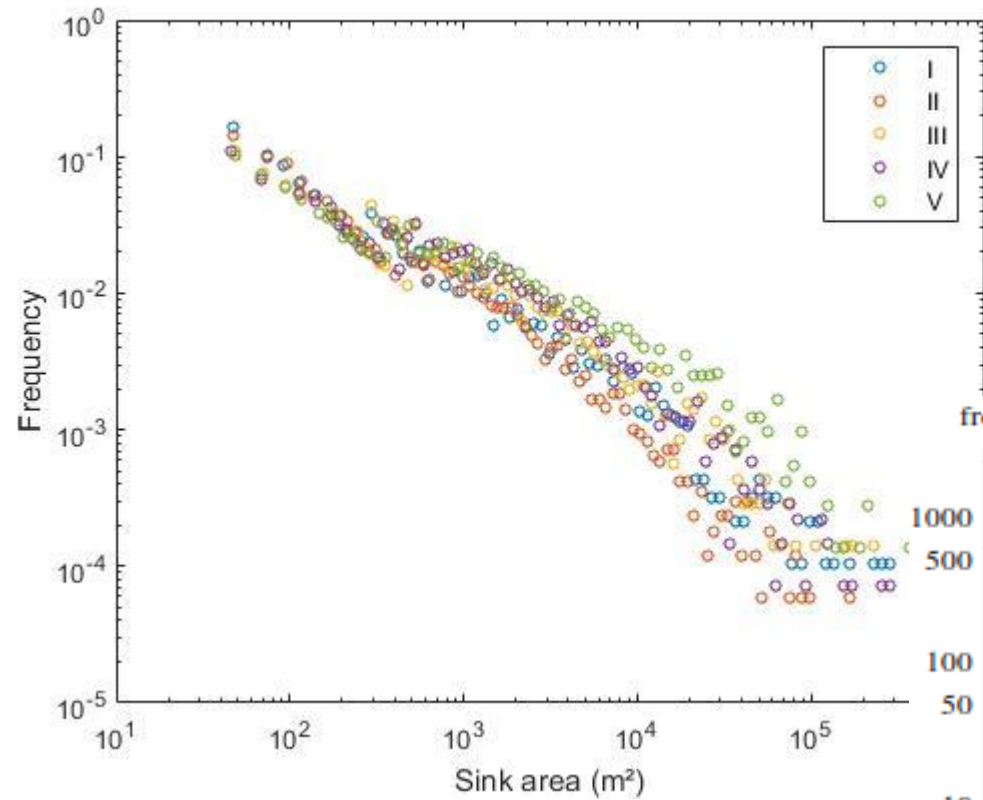
$$\frac{1}{2} (\partial_i v_j + \partial_j v_i) = A s^2 s_{ij}$$



Banerjee, Nainwal, Shankar, ...
(201*)



* evolution of surface topography in
Debris covered glaciers



Anderson, Scherler, Banerjee



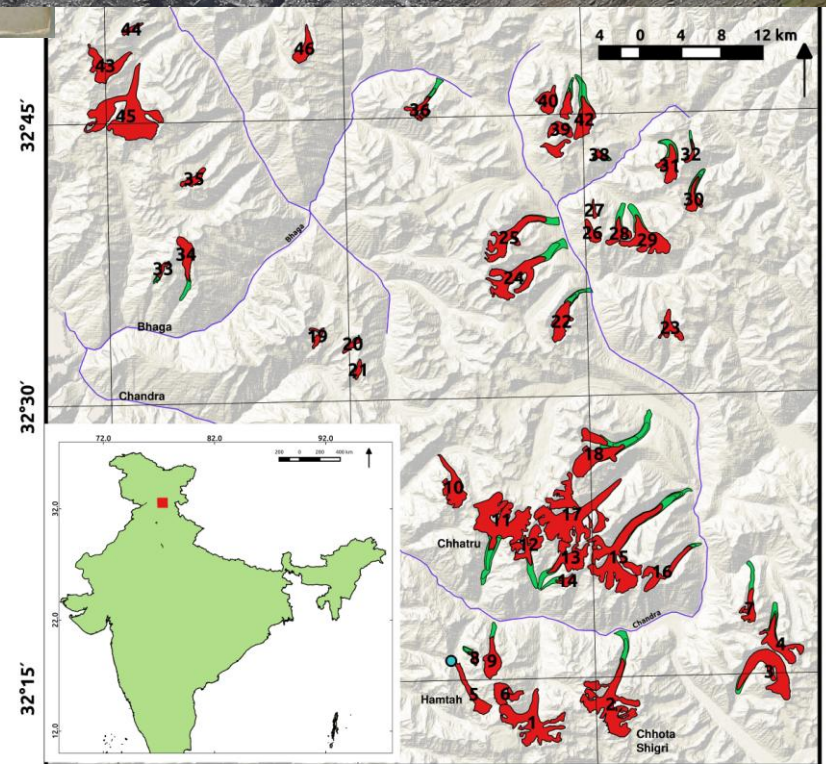
Kulu. Great Snowy Peak south of the Hamta Pass. 14,800 ft.

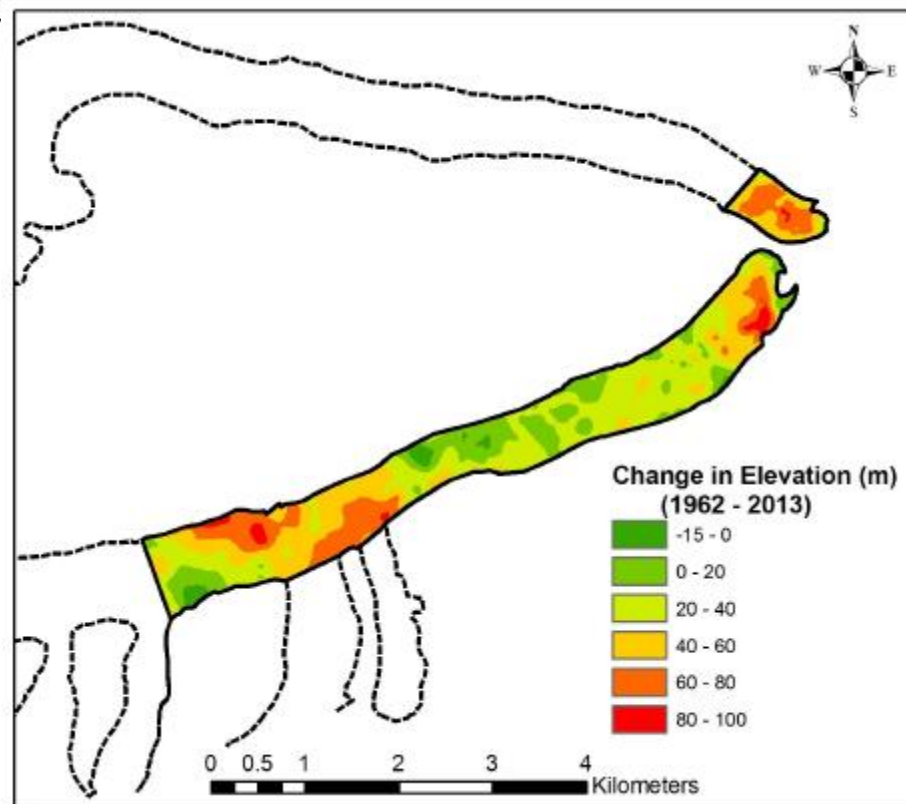
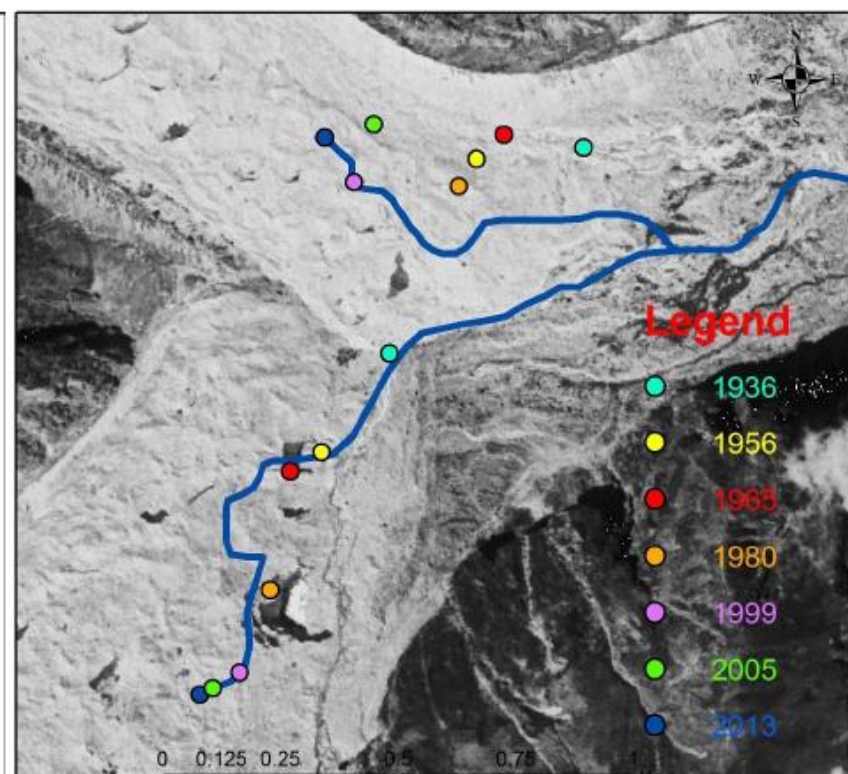
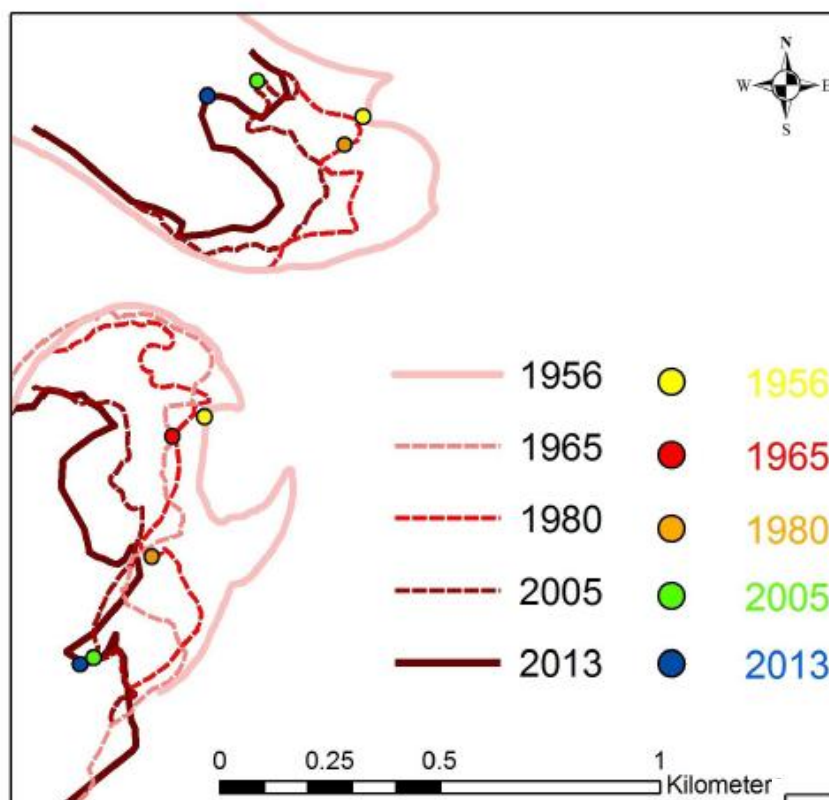


*Century-scale glacier fluctuaion in Lahaul
[Reshma]

* Debris-ice coupled flowline model
[Nikhil]

* Avalanche contribution in Hamtah
[Banerjee, Shankar, 2014]





[Nainwal, Banerjee and others, 2015]



Power-law profiles in great escarpment foothills

[Tejal, Reshma, Nishant, Sourav]

