

WORKSHOP ON DATA ASSIMILATION IN WEATHER AND CLIMATE MODELS 17 May 2024 International Centre for Theoretical Sciences, Tata Institute of Fundamental Research



Modelling Rainfall extremes and Floods over Karnataka

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Why Forecasting Extreme weather is so important for our Country?

10 countries hardest hit by weather disasters

- Top 10 countries with highest proportion of affected people over the total population (per 100,000 inhabitants)
- Top 10 countries with the highest absolute number of affected people (in million)



Forecast Leads- Applications





- In July 2021 alone, EREs caused property damage of around \$12 billion USD in Europe and China (Liang 2022)
- Economic loss of over \$3 billion USD per year in India (Swain et al. 2019)

It has changed in frequency and intensity as a result of climate change (Yin et al. 2022)

Damage Caused by Floods in India



Mohanty et al. 2020

Why Prediction of EREs is Important



weather events in India (Ray et al. 2021)

Lives Lost to Floods in India





Telemetric Rain-Gauge Stations (TRG) ~7000



Telemetric Weather Stations (TWS) ~950

(Temperature, Humidity, Wind Speed, Wind Direction)



Numerical models are not expected to be perfect: Forecast Models are susceptible to inherent errors Forecast errors can be due to imperfections in

- ✓ Data (Initial and Boundary conditions)
- ✓ Physics
- ✓ Numeric
- ✓ Approximation in basic equations
- ✓ Model grid

Systematic errors : Inherent with the model, generally follows specific pattern Relatively easier to eliminate How to eliminate: Using methods like bias correction

Non-Systematic errors: Random errors, not follows any specific pattern, Relatively difficult to eliminate, mostly due to errors in data (Initial and boundary condition) How to eliminate: Using methods like **Data Assimilation**



Model Configuration: Karnataka (3 Km); Bengaluru (666 m)



40 40° 20° 20° D4(0.666Km) D3d D2(6km) **0**° D1(18Km) 100° 60° 80 77.4 77.6 77.8 78.0 **(b)** 13.4 13.4 **Rural Bangalore** 13.2 13.2 13.0 13.0 **Ú**rban Bangalore 12.8 12.8 78.0 77.2 77.6 77.8 77.4

80°

100°

(a)

60°

- Domain Resolution: 27, 9 & 3 km for d01, d02 & d03 respectively
 Grid Points: 300 × 300 (d01), 352 × 352 (d02) and 271 × 376 (d03)
- Initial & Boundary Condition Data: NCEP Global Forecast System (GFS) 0.25° × 0.25° three hourly forecast data

Identification of EREs from coastal Karnataka

Distribution of rainfall events from coastal Karnataka region and ERE threshold at 99.993 percentile



Identification of EREs from South Interior Karnataka (SIK)

> Distribution of rainfall events from SIK region and ERE threshold at 99.993 percentile



Identification of EREs from North Interior Karnataka (NIK)

> Distribution of rainfall events from NIK region and ERE threshold at 99.993 percentile





Domain Configuration & Locations of EREs



> Domain Resolution:

27, 9 & 3km for d01, d02 & d03 respectively

→ Grid Points: $300 \times 300 (d01), 352 \times 352 (d02)$ and 271 × 376 (d03)

Initial & Boundary Condition Data:

NCEP Global Forecast System (GFS) $0.25^{\circ} \times 0.25^{\circ}$ three hourly forecast data

Spatial Distribution of Average Observed & Predicted Rainfall



Overprediction over the coastal region (windward side of western ghat (WG))
 Underprediction over leeward side of WG

Spatial Distribution of Errors in 24-h Accumulated Rainfall



- Bias observed in regions influenced by orographic lifting and associated convection
- Model underpredicted the rainfall over western ghats (WG)
- Model overpredicted (bias > 8mm) the high intense rainfall over the windward side of the mountainous WG

1200 UTC simulations shown lower RMSE than 0000 UTC simulations



source: Massimo Bonavita (ECMWF)

Data assimilation is the technique whereby observational data are combined with output from a numerical model to produce an optimal estimate of the evolving state of the system.

The Data Assimilation Process





Cost Function for Three Dimensional Variational (3DVar) DA

$$J = J_b + J_o$$

$$J = \frac{1}{2}(x - x_b)^T B^{-1}(x - x_b) + \frac{1}{2}(y_o - H(x))^T R^{-1}(y_o - H(x))$$

Background Term

Observation Term

Where, J_b – Background term J_o – Observation term

x – Analysis field value x_b – First guess

 y_o – Observed input H – Forward non-linear operator

B – Background Error Covariance Matrix

R – Observation Error Covariance Matrix

Optimal x_a is obtained by minimizing the cost function

 $\nabla_X J(X) = 0$

Source of the data	Variables assimilated in WRF
AIRS	Temperature, Relative Humidity, Dew Point Temperature
ASCAT	Wind Speed, Wind Direction
Buoy (RAMA & Moored)	Pressure, Temp, Relative Humidity, Wind Speed & Direction
KSNDMC	Temperature, Relative Humidity, Wind Speed
MODIS	Temperature, Dew Point Temperature
Radiosonde	Pressure, Height, Temperature, Dew Point Temperature, Relative Humidity, Wind Speed & Direction
SSMI	Wind Speed, Precipitable Water
WindSAT	Wind Speed, Wind Direction, & Precipitable Water



Locations of Observations Used in Assimilation



3D-Var DA in Cyclic mode



- Extreme rainfall event (June 21, 2015) simulated using 3D-Var in WRF
- 24 hour accumulated rainfall from June 21, 2015 0000 UTC June 22, 2015 0000 UTC
- Cyclic forecasts are updated with 06 hour forecast cycles
- Cyclic forecasts initiated at 0600 UTC on June 20, 2015
- Cyclic forecasts has reduced over-prediction

Comparison different assimilation Datasets



- Historical rainfall case over Bangalore: August 14, 2017
- Simulation performed using 3D-Var at 0000 UTC
- 24 hour accumulated rainfall is compared with ~6000 in-situ observations from TRG network of KSNDMC
- CNTL represents rainfall from experiment without assimilation
- DA_Satellite & in-situ obs represents rainfall from experiment with 3D-Var data assimilation using satellite profiles & in-situ data from buoy & station data from network of ~650 TWS stations
- DA_GTS_OBS represent rainfall from experiment with 3D-Var DA using NCEP prepbufr data

Spatial distribution of 24-h accumulated rainfall



- > In assimilation experiment using Ocean Winds, the spatial distribution of rainfall is simulated better.
- After assimilation, overprediction in comparison to the control experiment is reduced in the state's southeast regions.

Improvement Parameter (IP) in 24-h Accumulated Rainfall



$$IP = \left(1 - \frac{RMSE_{ASSI}}{RMSE_{CNT}}\right) \times 100$$

- Even after assimilation rainfall prediction is not improved over the WG region
- Maximum improvement : Ocean Winds
- Maximum Deterioration : Satellite Profile

Contingency table for skill score calculations

	Observation \geq Threshold	Observation < Threshold
$Forecast \ge Threshold$	a = Hits	b = False Alarms
Forecast < Threshold	c = Misses	d = Correct Negatives

Threat Score =
$$\frac{a}{a+b+c}$$

Improvement Parameter in skill score:

$$IP = \left(1 - \frac{|1 - SS_{ASSI}|}{|1 - SS_{CNT}|}\right) \times 100$$

Improvement Parameter based on Threat Score (TS)



- TS quantifies the skill of model in occurrence of the event of interest
- Assimilation with Ocean Winds
 & Station Data has shown maximum improvement
- Assimilation shown maximum deterioration over WG side with Station Data

Error in Categorical Rainfall



> Assimilation with Profile Data has improved area distribution of rainfall

IP in Wind Speed & Relative Humidity

IP in Wind Speed



Impact of assimilation

Seasonality in assimilation impact

Impact Ratio: value 0, nether improvement nor degradation values > 1, Improvement due to assimilation values < 1, Improvement due to assimilation



Urban ERE simulation: sub-kilometer resolution



MAP, Springer, 2021

Forecast Skill of Urban localized and Non-Localized EREs

$$RI_i = \left[\frac{I_i + L_i + A_i}{3}\right] * 100$$

Here, i stands for EREs (i=1,15)

Where, $I_{(i)}=\frac{1}{N}P(i,k)$ Quarterly journal of Royal Meteorological Society, 2018;143: 2340–2351

= 0, if relative error in magnitude of rainfall is above 50%, k = ERE number, k=1,34

$$L_{(i)} = \frac{1}{N} \sum_{k=1}^{N} P(i,k)$$

P(i,k) = 1, if rainfall is observed and forecasted over a hobli ,else, P(i,k) = 0 $A_{(i))=\frac{Nf(i)}{No(i)}}$

Where, Nf and No is the total number of hoblis where rainfall is forecasted and observed respectively



Cost Function for Hybrid-ETKF



Where,

- J_1 Traditional 3DVar background term associated with static covariance B
- A Defines spatial covariance of (a)
- ➤ Compared with a normal 3DVar, weighted sum of $J_1 \& J_e$ terms in Hybrid ETKF-3DVar cost function replaces the term J_b in 3DVar
- > To conserve the total background error variance $\frac{1}{\beta_1} + \frac{1}{\beta_2} = 1$

Flowchart of Hybrid ETKF-3DVar DA





Spatial Distribution of Average Rainfall



- > Intensity of rainfall is improved in HYBRID experiment
- ➢ Both experiment shows underprediction over WG

Spatial Distribution of Bias in 24-h Accumulated Rainfall



- Underprediction of rainfall is improved in HYBRID experiment compared to 3DVar
- The improvement in rainfall prediction in HYBRID is implied by a more number of TRG stations in the -2 to 2 bias group



> After the early hours, HYBRID simulations had less bias in rainfall prediction than 3DVar simulations.

Assimilation of GNSS Observations



Data and Model Domain



The location of (a) RO profiles (b) GNSS-GB (c) C2 profiles (d) Total number of GNSS observations used in study



WPS domain configuration of three nested domain d01 (27km), d02 (9km) and d03 (3km) with elevation

Experiment Design

- For initial and boundary conditions, NCEP-GFS dataset at 0.25 degree, starting 1200 UTC on the day prior the ERE day is used
- PREPBUFR: ship, radiosonde, aircraft, satellite profile etc.

Experiments	PREPBUFR	GNSS-GB	GPS-RO	C2
CNTL	\checkmark			
RO	\checkmark		\checkmark	
ZTD	\checkmark	\checkmark		
C2	\checkmark			\checkmark
RO+ZTD	\checkmark	\checkmark	\checkmark	
ZTD+C2	\checkmark	\checkmark		\checkmark
RO+ZTD+C2	\checkmark	\checkmark	\checkmark	\checkmark

Physics suite	Configuration	
Microphysics	New Thompson	
Radiation(Sw/Lw)	Dudhia/RRTM	
Land surface layer	Noah Land Surface	
PBL	Mellor-Yamada-Janjic	
Convection	Betts-Miller-Janjic	



Spatial distribution of 24 hour accumulated observed and modelled rainfall for 25102019.



Spatial distribution of RMSD of control experiment (mm) and improvement (%) from GNSS observations



Spatial distribution of ETS of control experiment, improvement (%) for GNSS observations and ETS for different rainfall threshold over Karnataka



Spatial distribution of RMSD and improvement parameters (%) for surface temperature at 00UTC



Vertical profile of RMSD (a) and improvement parameter (b) for relative humidity

INTEGRATED HEC RAS & SWAT MODEL FOR FLOOD MODELLING



Basin scale hydrological simulations: Quantification of spatio-temporal variability and change





1982 1983 1984 1985 1986 1987 1988 1989 1990 1991 1992 1993 1994 1995 1996 1997 1998 1999 2000 2001 2002 2003 2004 2005 2006 2007 2008 2009 2010 2011 2012 2013



Flood simulations:

- 1. Basin scale flood simulation (Periyar & Pamba basin)
- 2. Flood simulation in paddy fields (Kuttanadu)







Storm Water Management Model (SWMM) simulation for sub-catchment, Node flooding and link flow over **Koramangala, Challaghatta watershed valley** as a result of extreme rainfall events on **7**th **September 2015**. (KSNDMC Rainfall data)



